

# Vertical Distribution of Soil Carbon, Nitrogen, and Phosphorus in Typical Chinese Terrestrial Ecosystems

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**Abstract:** Characterization of the vertical distribution of soil organic carbon (C), nitrogen (N), and phosphorus (P) may improve our ability to accurately estimate soil C, N, and P storage. Based on a database of 21 354 records in 74 long-term monitoring plots from 2004 to 2013 in the Chinese Ecosystem Research Network (CERN), we built fitting functions to quantify the vertical distribution of soil C, N, and P (up to 100 cm depth) in the typical Chinese terrestrial ecosystems. The decrease of soil C, N, and P content with depth can be well fitted with various mathematical functions. The fitting functions differed greatly between artificial (agriculture) and natural (desert, forest, and grassland) ecosystems, and also differed with respect to soil C, N, and P content. In both the artificial and natural ecosystems, the best fitting functions were exponential functions for C, quadratic functions for N, and quadratic functions for P. Furthermore, the stoichiometric ratios of soil C, N, and P were ranked in descending order: grassland > forest > agriculture > desert, and were also associated with climate. This study is the first to build the fitting functions for the profile distribution of soil C, N, and P in China at a national scale. Our findings provide a scientific basis to accurately assess the storage of C, N, and P in soils at a large scale, especially for the integrative analysis of historical data.

**Keywords:** soil profile; storage; stoichiometry; vertical distribution; China

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## 1 Introduction

Soil organic carbon (SOC), nitrogen (N), and phosphorus (P) are important indicators for soil fertility and quality (Andrews *et al.*, 2004; Lal, 2004; Schindler, 2006). Soil is the largest SOC pool of terrestrial ecosystems (1500 Pg C) (Jobbagy and Jackson, 2000). Many studies have demonstrated that terrestrial ecosystems have a tremendous capacity to sequester atmospheric CO<sub>2</sub>; thus, rational management could help partially offset anthropogenic CO<sub>2</sub> increase (Conant *et al.*, 2001;

He *et al.*, 2008). N is essential for plant growth, and it is estimated that vegetation adsorbs about 50%–70% N from the soil (Dewes, 1999). Thus, N addition may significantly enhance primary productivity in most terrestrial ecosystems. Similarly, P is another important element for plant growth (Compton *et al.*, 2000), and is mainly derived from soil parent material and fertilization (Wang *et al.*, 1999). Consequently, P is considered as a limiting nutrient for plant growth at the geological time scale (Compton *et al.*, 2000).

The cycles of C, N, and P in soils are closely corre-

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lated (Agren *et al.*, 1991; Vitousek, 2004). Global warming resulting from an increase in atmospheric CO<sub>2</sub> alters the plant growth environment, and affects N and P turnover (Robinson *et al.*, 1995). For example, rising temperature is expected to promote soil N mineralization and to enhance plant growth (Tong *et al.*, 2005). Available P is influenced by the parent rock over a long time, but is mainly derived from the decomposition of litter and soil organic matter (SOM) over a short time (Kellogg and Bridgham, 2003). Climate warming impacts soil P availability by altering temperature and moisture, which control the decomposition of litter, SOM, and soil parent material (Robinson *et al.*, 1995). Overall, the ecosystem C budget is restricted by soil N and P availability (Reich and Oleksyn, 2004), because they affect plant growth and control the accumulation and decomposition of SOM (Hobbie and Vitousek, 2000; Harrington *et al.*, 2001). Furthermore, soil N and P availability may affect the N and P content of plants and thus plant photosynthesis (Cordell *et al.*, 2001a; 2001b). Understanding the profile distributions of soil C, N, and P may therefore provide a solid foundation for exploring how these cycles are coupled.

Some studies have demonstrated that the vertical distribution of soil C, N, and P at specific sites or regions. Alvarez *et al.* (2011) studied the vertical distribution of C and N in the 0–30 cm soil layer under different management practices in Argentina. van der Wal *et al.* (2007) reported the P vertical distribution of 0–30 cm soil at an abandoned farmland in the Netherlands. Furthermore, Liu *et al.* (2012) reported the C vertical distribution of brown soil to a depth of 60 cm under different land use types in China. Qiu *et al.* (2012) investigated the profile distribution of P in red, alluvial, and purple soils in China. Wang *et al.* (2004) estimated the SOC vertical distribution and its relationship to climate and vegetation in Chinese terrestrial ecosystems based on soil geographic and vertical attribute data from the Chinese Second Soil Survey Database. However, it remains unclear whether the vertical distributions of soil C, N, and P can be fitted by mathematic functions, and whether the fitting functions remain consistent in different ecosystem types.

To date, the fitting functions of the vertical distributions of soil C, N, and P have not been reported at large scale. Possible reasons included the inconsistency of

sampling and analytical methods in historical datasets, a deficiency of replicates, and a lack of data for deeper soil layers. Fortunately, the long-term monitoring program of the Chinese Ecosystem Research network (CERN) has overcome these limitations, through implementing consistent sampling designs and measurement methods over a long time period. In this study, a database of containing monitoring records at 74 long-term experimental plots in CERN were used to quantitatively assess the vertical distributions of C, N, and P in the 0–100 cm soil profile. The main objectives were to: 1) establish the fitting functions for the vertical distribution of soil C, N, and P in Chinese typical terrestrial ecosystems; 2) explore whether the fitting functions are consistent among different ecosystem types, and 3) provide scientific parameters for the accurate estimation of soil C and N storage at regional and national scales.

## 2 Materials and Methods

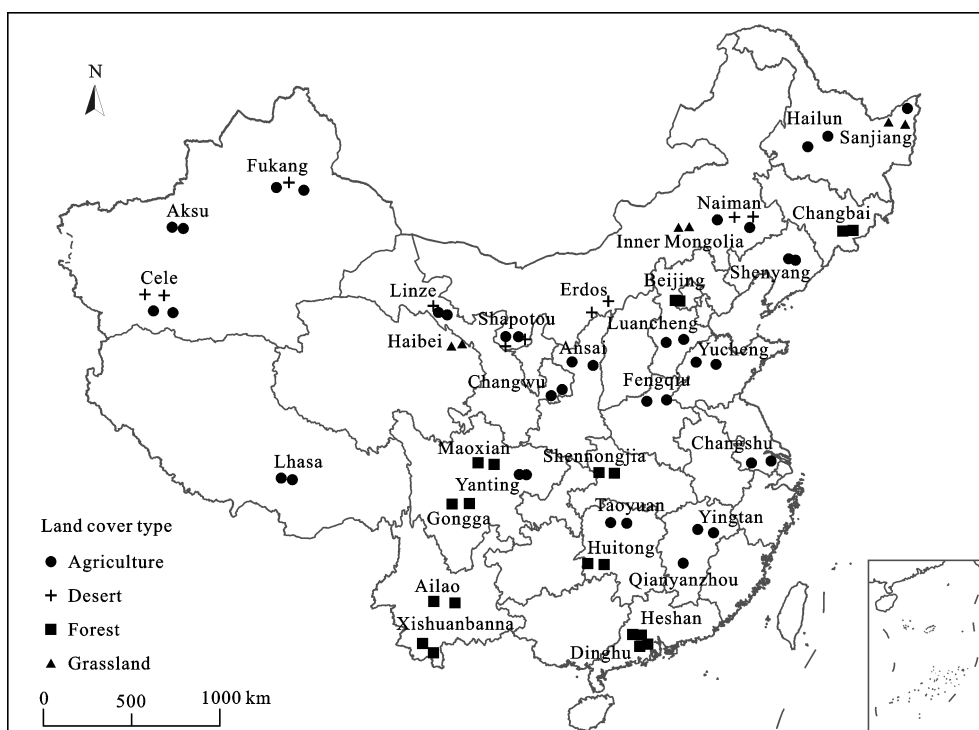
### 2.1 Study area and data processing

The study area covers most of the typical terrestrial ecosystems in China. In order to explore the fitting functions of soil C, N, and P profile distribution in typical terrestrial ecosystems in China, we used the data from the long-term monitoring experimental plots at field stations of CERN ([www.cerndata.ac.cn](http://www.cerndata.ac.cn)) (Fu *et al.*, 2010) (Fig. 1).

In the data processing, we only selected 33 field stations that have been regularly monitored for more than 10 years. At each station, more than 2 plots were selected. If a specific station contained different ecosystem types (agriculture, forest, grassland, and desert), we selected 2 plots for each ecosystem type. A total of 74 experimental plots were selected, and the description (e.g., ecosystem type, topography, soil type, and dominant species) for each plot was provided in Table 1.

### 2.2 Measurements of C, N, and P

A total of 21 354 records were available for soil C, N, and P in 0–100 cm soil depth from 2004 to 2013. The SOC content (g/kg) was measured using the modified Mebius method (Nelson and Sommers, 1982). Total N (TN) content (g/kg) was measured using the modified Kjeldahl wet digestion procedure (Gallaher *et al.*, 1976). Total P (TP) content (g/kg) was determined by the ammonium molybdate method after persulfate oxidation (Kuo, 1996).



**Fig. 1** Long-term experimental plots of Chinese Ecosystem Research Network (CERN) used in this study. Different ecosystem types of experimental plots in each ecological station were represented as different shape in legend. Names were informal abbreviations for each station

**Table 1** Information of long-term monitoring plots of Chinese Ecosystem Research Network (CERN) in this study

Ecological station type	Station	Ecosystem type	Plot No.	Topography	Soil type	Dominant species
Agriculture	Aksu	Agriculture	A	Flat ground	Irrigation-farming soil	Cotton
			B	Flat ground	Irrigation-farming soil	Cotton
	Ansai	Agriculture	A	Terrace	Cultivated loessial soil	Spring maize, maize, soybean
			B	Terrace	Cultivated loessial soil	Spring maize, maize, soybean
	Changshu	Agriculture	A	Flat ground	Paddy soil	Wheat, rice
			B	Flat ground	Paddy soil	Wheat, rice
	Fengqiu	Agriculture	A	Flat ground	Fluvo-aquic soil	Maize, wheat
			B	Flat ground	Fluvo-aquic soil	Maize, wheat
	Hailun	Agriculture	A	Flat ground	Black soil	Maize, wheat, soybean
			B	Flat ground	Black soil	Maize, soybean
	Lhasa	Agriculture	A	Flat ground	Fluvo-aquic soil	Winter wheat, Spring barley
			B	Flat ground	Fluvo-aquic soil	Winter wheat, rape
	Linze	Desert	A	Flat ground	Gray-brown desert soil	<i>Nitraria sphaerocarpa</i>
		Agriculture	A	Flat ground	Aeolian sandy soil	Maize, wheat, spring wheat
			B	Flat ground	Aeolian sandy soil	Maize, wheat, alfalfa
	Luancheng	Agriculture	A	Flat ground	Meadow cinnamon soil	Wheat, maize
			B	Flat ground	Meadow cinnamon soil	Wheat, maize
	Qianyanzhou	Agriculture	A	Flat ground	Hydromorphic paddy soil	Rice
	Shenyang	Agriculture	A	Flat ground	Aquic brown soil	Maize, soybean
			B	Flat ground	Aquic brown soil	Maize, soybean
	Taoyuan	Agriculture	A	Flat ground	Paddy soil	Rice
			B	Flat ground	Paddy soil	Rice
	Yanting	Agriculture	A	Slop land	Calcareous purple soil	Maize, wheat
			B	Slop land	Calcareous purple soil	Maize, wheat, rape
	Yingtian	Agriculture	A	Gentle slope	Red soil	Peanut

Continued table

Ecological station type	Station	Ecosystem type	Plot No.	Topography	Soil type	Dominant species
Agriculture	Yucheng	Agriculture	B	Flat ground	Red soil	Peanut
			A	Flat ground	Saline flavo-aquic soil	Maize, wheat
	Changwu	Agriculture	B	Flat ground	Saline flavo-aquic soil	Maize, wheat
			A	Flat ground	Dark loessial soil	Wheat, maize
			B	Flat ground	Dark loessial soil	Wheat
Forest	Ailao	Forest	A	Slope land	Mountain yellow-brown soil	Evergreen broad-leaved forest
			B	Slope land	Brown soil	<i>Planus coppice</i>
	Xishuanbanna	Forest	A	Slope land	Humid-thermo ferralitic soil	Seasonal rainforest
			B	Slope land	Humid-thermo ferralitic soil	Secondary forest (valley)
	Heshan	Forest	A	Low hill	Lateritic red soil	Acacia mangium forest
			B	Low hill	Lateritic red soil	<i>Schima superba</i> , <i>Castanopsis hystrix</i>
	Beijing	Forest	A	Slope land	Mountain brown soil	Deciduous broadleaved mixed forest
			B	Slope land	Mountain brown soil	North China larch
	Dinghu	Forest	A	Slope land	Lateritic red soil	Subtropical evergreen broadleaved forest
			B	Slope land	Lateritic red soil	Masson pine forest
	Gongga	Forest	A	Slope land	Brown coniferous forest soil	<i>Abies fabri</i>
			B	Gentle slope	Skeleton soil	<i>A. fabri</i> , <i>Populus purdomii</i>
	Huitong	Forest	A	Slope land	Yellow soil	Coniferous forest
			B	Slope land	Yellow soil	Coniferous forest
	Maoxian	Forest	A	Slope land	Cinnamon soil	Coniferous and broadleaved mixed forest
			B	Slope land	Brown soil	Deciduous broad-leaved shrub
	Shennongjia	Forest	A	Slope land	Mountain yellow-brown soil	Evergreen broadleaf mixed forest
			B	Slope land	Mountain dark-brown soil	Coniferous forest
	Changbai	Forest	A	Flat ground	Dark-brown soil	Broad-leaved korean pine forest
			B	Flat ground	Dark-brown soil	Birch forest
Grassland	Inner Mongolia	Grassland	A	Flat ground	Dark chestnut soil	<i>Leymus chinensis</i>
			B	Flat ground	Chestnut soil	<i>Stipa grandis</i>
	Haibei	Grassland	A	Flat ground	Felty soil	Alpine <i>kobresia</i> meadow
			B	Flat ground	Dark felty soil	Alpine <i>kobresia</i> meadow
	Sanjiang	Grassland	A	Depression	Bog soil	Carex
			B	Flat ground	Bog soil	<i>Deyeuxia angustifolia</i>
	Agriculture	A	Flat ground	Albic soil	Soybean, wheat, maize	
Desert	Erdos	Desert	A	Gentie dune	Aeolian sandy soil	<i>Artemisia ordosica</i>
			B	Gentie dune	Aeolian sandy soil	<i>A. ordosica</i>
	Cele	Desert	A	Dune	Aeolian sandy soil	<i>Alhagi sparsifolia</i>
			B	Dune	Aeolian sandy soil	<i>A. sparsifolia</i>
		Agriculture	A	Flat ground	Aeolian sandy soil	Cotton
			B	Flat ground	Aeolian sandy soil	Cotton
	Fukang	Desert	A	Flat ground	Aeolian sandy soil	<i>Haloxylon ammodendron</i>
		Agriculture	A	Flat ground	Gray desert soil	Cotton, wheat, maize
			B	Flat ground	Gray desert soil	Wheat, winter wheat
			Naiman	Desert	A	Flat ground
	B	Gentle slope		Aeolian sandy soil	<i>Caragana microphylla</i>	
		Agriculture	A	Flat ground	Meadow aeolian sandy soil	Soybean, wheat, maize
			B	Flat ground	Meadow aeolian sandy soil	Maize
	Shapotou	Desert	A	Flat ground	Aeolian sandy soil	<i>Artemisia ordosica</i>
B			Fat ground	Aeolian sandy soil	<i>Artemisia ordosica</i>	
Agriculture		A	Flat ground	Cumulated irrigated soil	Maize, wheat	
		B	Flat ground	Cumulated irrigated soil	Maize, wheat	

Note: A and B represent different experimental plots with different dominant species or ecosystem types in the same ecological station

### 2.3 Classification of ecosystem types and climate zones

All plots were divided into four ecosystem types (agriculture, forest, grassland, and desert) and six climatic zones (temperate, warm temperate, north subtropical, mid-subtropical, south subtropics, and Tibetan Plateau zones) (Zhang *et al.*, 2010) to explore the general trends.

### 2.4 Curve estimation

The soil C, N, and P content was first averaged at different depths for the 10-year monitoring period. The data were then analyzed by Curve Estimation of SPSS 13.0, which contain quadratic, compound, growth, logarithmic, S, exponential, inverse, and power functions.

When the coefficient ( $R^2$ ) of the fitting functions was similar, Akaike Information Criterion (AIC) and Bayesian Information Criterion (BIC) were used to determine the best fit function, from which functions with smaller AIC and BIC values were selected (Bozdogan, 1987; Aho *et al.*, 2014). In practice, AIC and BIC were calculated as follows:

$$AIC = n \times \ln(Rss/n) + 2(k+1) \quad (1)$$

$$BIC = n \times \ln(Rss/n) + 2(k+1) \times \ln(n) \quad (2)$$

where  $RSS$  is the residual sum of squares;  $n$  is the number of observations, and  $k$  is the number of estimated parameters.

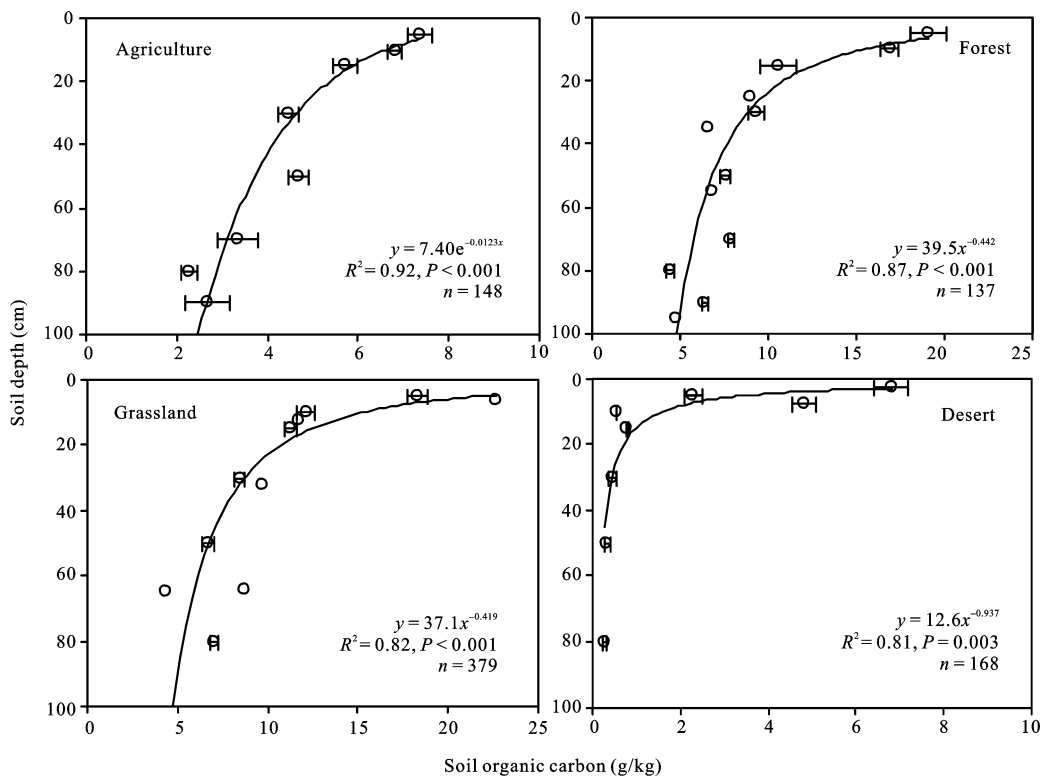
## 3 Results

### 3.1 Vertical distribution of soil C, N, and P content

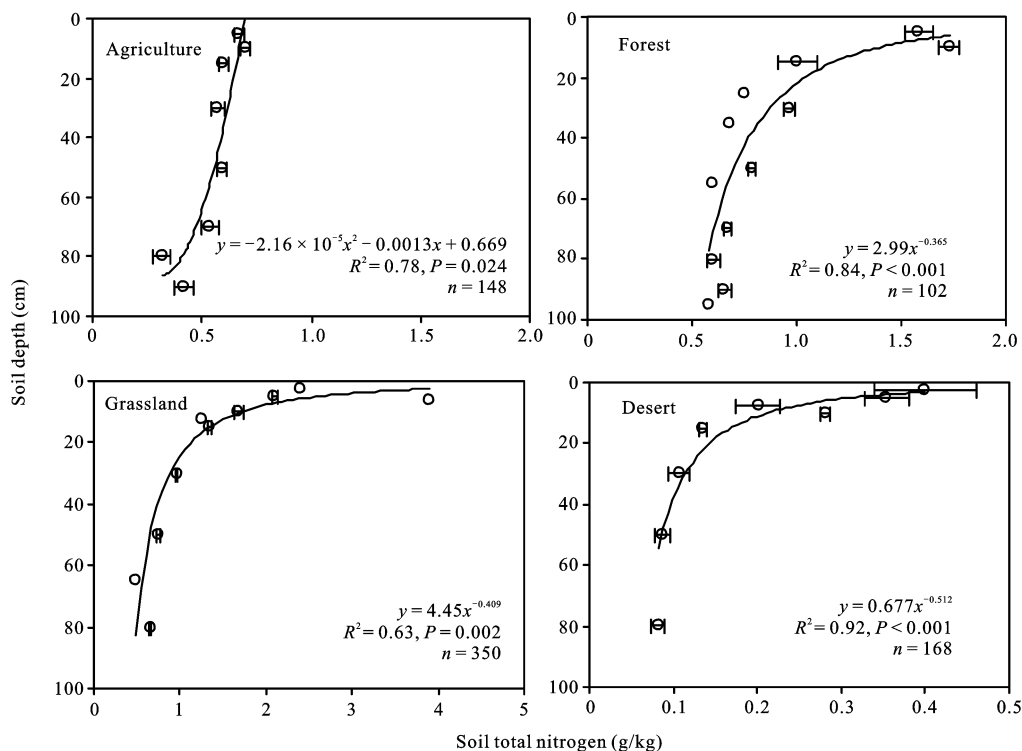
The profile distribution of soil C, N, and P content was higher in the upper soil layer compared to the deep soil layer, irrespective of vegetation types or climate zones. For example, Figs. 2–4 show that soil C, N, and P content significantly decreases with increasing soil depth in agriculture (Fengqiu-A), forest (Xishuangbanna-A), grassland (Inner Mongolia-A), and desert ecosystems (Shapotou-A). Similar trends were observed in other plots (data not shown).

### 3.2 Fitting functions for vertical distribution of soil C, N, and P

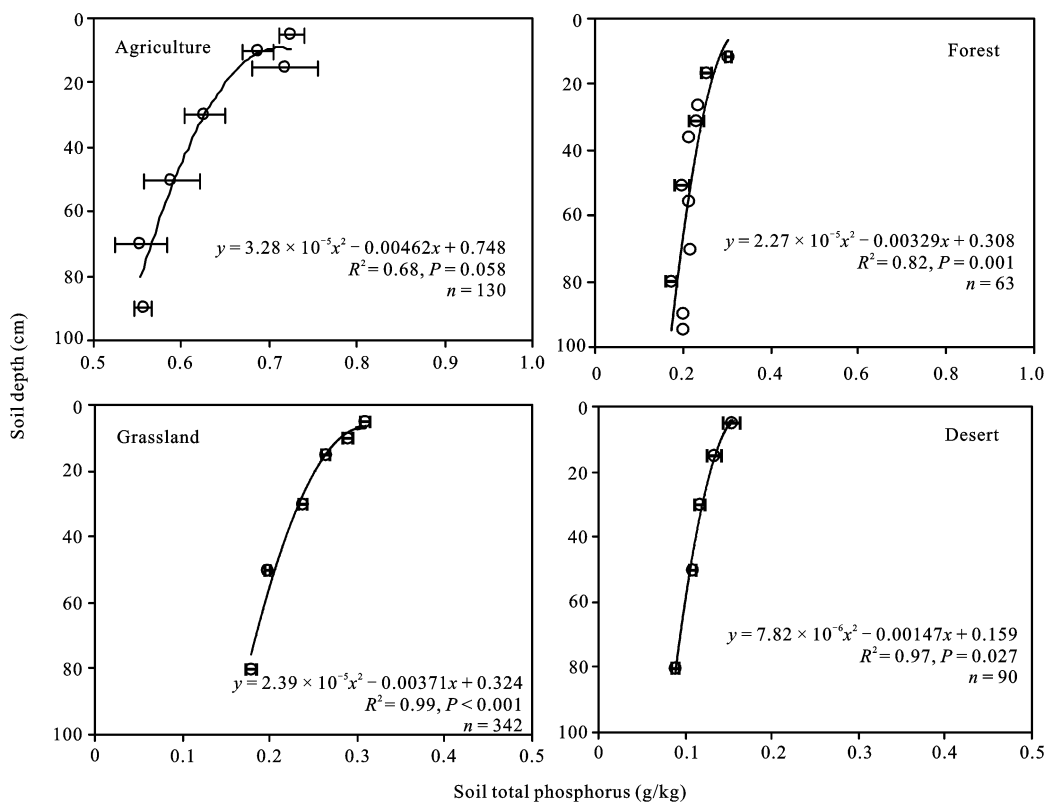
As shown in Table 2, the vertical distribution of soil C,



**Fig. 2** Vertical distribution of soil organic carbon of typical terrestrial ecosystems. Agriculture: Fengqiu-A; forest: Xishuangbanna-A; grassland: Inner Mongolia-A; desert: Shapotou-A. In the functions,  $y$  and  $x$  represent soil organic carbon (g/kg) and soil depth (cm), respectively



**Fig. 3** Vertical distribution of soil total nitrogen of typical terrestrial ecosystems. Agriculture: Fengqiu-A; forest: Xishuangbanna-A; grassland: Inner Mongolia-A; desert: Shapotou-A. In the functions,  $y$  and  $x$  represent soil total nitrogen (g/kg) and soil depth (cm)



**Fig. 4** Vertical distribution of soil total phosphorus of typical terrestrial ecosystems. Agriculture: Fengqiu-A; forest: Xishuangbanna-A; grassland: Inner Mongolia-A; desert: Shapotou-A. In the functions,  $y$  and  $x$  represent soil total phosphorus (g/kg) and soil depth (cm)

**Table 2** Fitting functions of soil organic carbon (SOC), total nitrogen (TN), and total phosphorus (TP) against soil depth in 0–100 cm soil profile of typical terrestrial ecosystems in China

Ecosystem type	Station	Plot No.	SOC			TN			TP		
			Function	$R^2$	$P$	Function	$R^2$	$P$	Function	$R^2$	$P$
Agriculture	Aksu	A	$y=6.26e^{-0.0127x}$	0.94	0.006	$y=7.41 \times 10^{-5}x^2-0.0124x+0.740$	0.92	0.085	$y=8.07 \times 10^{-5}x^2-0.0111x+0.926$	0.85	0.150
		B	$y=4.15e^{-0.00855x}$	0.83	0.033	$y=7.25 \times 10^{-5}x^2-0.0095x+0.514$	0.95	0.054	$y=1.57 \times 10^{-5}x^2-0.0043x+0.818$	0.95	0.049
	Ansai	A	$y=6.28e^{-0.0137x}$	0.96	0.004	$y=5.04 \times 10^{-5}x^2-0.0095x+0.710$	0.99	0.005	$y=2.93 \times 10^{-5}x^2-0.0046x+0.731$	0.99	0.006
		B	$y=4.91e^{-0.00768x}$	0.99	0.001	$y=-1.07 \times 10^{-6}x^2-0.0021x+0.495$	0.99	0.008	$y=8.57 \times 10^{-6}x^2-0.0022x+0.668$	0.89	0.113
	Changshu	A	$y=26.8e^{-0.0361x}$	0.78	<0.001	$y=4.35 \times 10^{-4}x^2-0.0649x+2.78$	0.95	<0.001	$y=-3.66 \times 10^{-6}x^2-0.0025x+0.819$	0.47	0.057
		B	$y=22.7e^{-0.0213x}$	0.94	<0.001	$y=4.12 \times 10^{-4}x^2-0.0574x+2.51$	0.95	<0.001	$y=-9.48 \times 10^{-5}x^2-0.0022x+0.734$	0.91	0.002
	Fengqu	A	$y=7.40e^{-0.0123x}$	0.92	<0.001	$y=-2.16 \times 10^{-5}x^2-0.0013x+0.669$	0.78	0.024	$y=3.28 \times 10^{-5}x^2-0.0046x+0.748$	0.68	0.058
		B	$y=6.40e^{-0.00942x}$	0.90	<0.001	$y=-2.10 \times 10^{-5}x^2-0.0013x+0.584$	0.78	0.024	$y=3.10 \times 10^{-5}x^2-0.0040x+0.677$	0.89	0.004
	Hailun	A	$y=29.7e^{-0.0151x}$	0.97	<0.001	$y=2.78e^{-0.0169x}$	0.94	0.001	$y=1.22 \times 10^{-4}x^2-0.0153x+1.03$	0.89	0.036
		B	$y=29.8e^{-0.0155x}$	0.97	<0.001	$y=2.60e^{-0.0161x}$	0.96	0.001	$y=4.30 \times 10^{-5}x^2-0.0085x+0.920$	0.94	0.015
	Lhasa	A	$y=15.3e^{-0.0221x}$	0.98	0.001	$y=1.61e^{-0.0185x}$	0.98	0.002	$y=6.57 \times 10^{-5}x^2-0.0070x+0.858$	0.89	0.327
		B	$y=10.6e^{-0.0199x}$	0.99	<0.001	$y=1.17 \times 10^{-4}x^2-0.0174x+1.11$	0.97	0.032	$y=-6.91 \times 10^{-6}x^2-0.0034x+0.782$	0.99	0.066
	Linze	A	$y=28.9x^{-0.891}$	0.91	0.003	$y=1.54 \times 10^{-4}x^2-0.0190x+0.643$	0.95	0.010	$y=6.09 \times 10^{-5}x^2-0.0072x+0.445$	0.96	0.007
		B	$y=59.6x^{-1.06}$	0.93	0.002	$y=2.41 \times 10^{-4}x^2-0.0297x+0.919$	0.98	0.004	$y=7.87 \times 10^{-5}x^2-0.0099x+0.537$	0.95	0.012
	Luancheng	A	$y=30.2x^{-0.538}$	0.92	0.003	$y=2.77 \times 10^{-4}x^2-0.0342x+1.42$	0.98	0.003	$y=-0.175 \ln x+1.10$	0.86	0.008
		B	$y=20.1x^{-0.409}$	0.86	0.007	$y=2.10 \times 10^{-4}x^2-0.0264x+1.19$	0.94	0.016	$y=1.03 \times 10^{-4}x^2-0.0152x+0.883$	0.99	0.009
	Qianyanzhou	A	$y=57.8x^{-0.780}$	0.91	<0.001	$y=4.23 \times 10^{-4}x^2-0.0482x+1.56$	0.92	<0.001	$y=4.84 \times 10^{-5}x^2-0.0058x+0.291$	0.94	0.003
		A	$y=16.3x^{-0.243}$	0.90	0.001	$y=-0.154 \ln x+1.39$	0.83	0.004	$y=6.37 \times 10^{-5}x^2-0.0058x+0.582$	0.87	0.047
	Shenyang	A	$y=17.2x^{-0.243}$	0.78	0.009	$y=-0.130 \ln x+1.31$	0.76	0.011	$y=2.65 \times 10^{-6}x^2+0.0016x+0.477$	0.40	0.362
		B	$y=23.3x^{-0.311}$	0.82	0.005	$y=1.35e^{-0.00916x}$	0.73	0.014	$y=6.93 \times 10^{-5}x^2-0.0080x+0.611$	0.90	0.011
	Tao Yuan	A	$y=18.1e^{-0.0165x}$	0.97	<0.001	$y=2.17 \times 10^{-4}x^2-0.0339x+1.83$	0.95	0.013	$y=3.22 \times 10^{-5}x^2-0.0036x+0.505$	0.99	0.001
		B	$y=11.9x^{-0.332}$	0.94	0.006	$y=-0.172 \ln x+1.12$	0.89	0.017	$y=9.57 \times 10^{-5}x^2-0.0091x+0.750$	0.69	0.314
	Yanting	A	$y=6.14e^{-0.0121x}$	0.57	0.049	$y=-1.13 \times 10^{-4}x^2+6.41 \times 10^{-5}x+0.720$	0.07	0.867	$y=-1.08 \times 10^{-4}x^2-0.0053x+0.670$	0.12	0.776
		B	$y=6.50e^{-0.0193x}$	0.94	<0.001	$y=5.95 \times 10^{-5}x^2-0.0101x+0.726$	0.94	0.004	$y=5.82 \times 10^{-5}x^2-0.0074x+0.492$	0.95	0.003
	Yingtian	A	$y=8.51e^{-0.0204x}$	0.98	<0.001	$y=1.21 \times 10^{-4}x^2-0.0171x+0.925$	0.97	0.001	$y=1.56 \times 10^{-4}x^2-0.0179x+0.751$	0.98	<0.001
		B	$y=31.5x^{-0.609}$	0.91	0.003	$y=1.92 \times 10^{-4}x^2-0.0263x+1.18$	0.93	0.017	$y=9.45 \times 10^{-5}x^2-0.0127x+0.989$	0.89	0.038
	Yucheng	A	$y=5.85e^{-0.0127x}$	0.98	<0.001	$y=4.43 \times 10^{-5}x^2-0.0080x+0.693$	0.99	0.002	$y=9.20 \times 10^{-6}x^2-0.0017x+0.661$	0.88	0.044
		B	$y=10.6x^{-0.139}$	0.79	0.001	$y=4.36 \times 10^{-5}x^2-0.0075x+0.991$	0.87	0.001	$y=6.12 \times 10^{-5}x^2-0.0083x+0.896$	0.85	0.001
	Changwu	A	$y=10.4x^{-0.141}$	0.76	0.024	$y=7.46 \times 10^{-5}x^2-0.0098x+0.943$	0.91	0.027	$y=4.63 \times 10^{-5}x^2-0.0069x+0.848$	0.92	0.024
		B	$y=2.90e^{-0.00537x}$	0.83	0.004	$y=1.07 \times 10^{-5}x^2-0.0027x+0.429$	0.80	0.041	$y=2.92 \times 10^{-5}x^2-0.0034x+0.641$	0.98	0.004
	Cele	A	$y=6.84x^{-0.292}$	0.73	0.015	$y=5.48 \times 10^{-5}x^2-0.0081x+0.564$	0.86	0.018	$y=5.60 \times 10^{-5}x^2-0.0073x+0.795$	0.86	0.052
		B	$y=7.24e^{-0.0188x}$	0.98	<0.001	$y=5.16 \times 10^{-5}x^2-0.0109x+0.742$	0.93	0.001	$y=-2.21 \times 10^{-5}x^2+5.88 \times 10^{-4}x+0.538$	0.02	0.951
	Fukang	A	$y=10.3x^{-0.273}$	0.78	0.020	$y=1.48 \times 10^{-4}x^2-0.0127x+0.587$	0.87	0.045	$y=-1.14 \times 10^{-4}x^2-0.0147x+0.510$	0.18	0.743
		B	$y=8.41x^{-0.192}$	0.42	0.002	$y=1.47/x+0.443$	0.45	0.001	$y=3.51 \times 10^{-5}x^2-0.0081x+0.741$	0.36	0.019
	Naiman	A	$y=e^{3.59/x+0.873}$	0.85	0.009	$y=1.04 \times 10^{-4}x^2-0.0109x+0.549$	0.93	0.019	$y=1.73 \times 10^{-4}x^2-0.0189x+0.940$	0.42	0.438
		B	$y=6.93e^{-0.0309x}$	0.89	<0.001	$y=2.08 \times 10^{-4}x^2-0.0255x+0.852$	0.97	<0.001	$y=1.22 \times 10^{-4}x^2-0.0158x+0.637$	0.87	0.001
	Shapotou	A	$y=40.6x^{-0.954}$	0.89	<0.001	$y=2.08 \times 10^{-4}x^2-0.0256x+0.832$	0.93	0.001	$y=1.24 \times 10^{-4}x^2-0.0153x+0.553$	0.95	0.051
		B	$y=42.9x^{-0.376}$	0.64	0.001	$y=6.49x^{-0.538}$	0.71	<0.001	$y=1.94 \times 10^{-5}x^2-0.0089x+0.969$	0.39	0.065

Continued table

Ecosystem type	Station	Plot No.	SOC			TN			TP		
			Function	$R^2$	$P$	Function	$R^2$	$P$	Function	$R^2$	$P$
Forest	Ailao	A	$y=132e^{-0.0199x}$	0.99	<0.001	$y=8.90e^{-0.0186x}$	0.99	<0.001	$y=1.26e^{-0.00695x}$	0.96	<0.001
		B	$y=119e^{-0.0279x}$	0.98	<0.001	$y=6.08e^{-0.0260x}$	0.99	<0.001	$y=8.96 \times 10^{-5}x^2-0.0122x+0.822$	0.50	0.355
	Beijing	A	$y=21.0e^{-0.0299x}$	0.98	<0.001	$y=0.00109x^2-0.133x+4.41$	0.99	<0.001	$y=1.37 \times 10^{-4}x^2-0.0148x+0.575$	0.98	0.004
		B	$y=25.8x^{-0.0325}$	0.99	<0.001	$y=4.54e^{-0.0309x}$	0.99	<0.001	$y=5.63 \times 10^{-5}x^2-0.0084x+0.552$	0.99	<0.001
	Dinghu	A	$y=93.3x^{-0.653}$	0.87	<0.001	$y=3.68 \times 10^{-4}x^2-0.0426x+1.70$	0.53	0.033	$y=0.676x+0.174$	0.88	0.001
		B	$y=37.1x^{-0.611}$	0.77	<0.001	$y=2.35x+0.248$	0.81	<0.001	$y=5.04 \times 10^{-5}x^2-0.0029x+0.225$	0.55	0.060
	Gongga	A	$y=481x^{-0.919}$	0.62	<0.001	$y=23.5x^{-0.697}$	0.72	0.001	$y=e^{-4.22/x+0.376}$	0.93	<0.001
		B	$y=e^{11.6x+1.59}$	0.98	<0.001	$y=58.1x^{-1.43}$	0.91	<0.001	$y=1.08x+0.977$	0.42	0.240
	Heshan	A	$y=66.6x^{-0.652}$	0.79	0.001	$y=7.36x+0.396$	0.91	<0.001	$y=3.76 \times 10^{-9}x^2+6.47 \times 10^{-4}x+0.212$	0.77	0.026
		B	$y=17.8e^{-0.0235x}$	0.84	0.001	$y=-0.364 \ln x+2.05$	0.98	<0.001	$y=1.56 \times 10^{-5}x^2+1.26 \times 10^{-4}x+0.284$	0.18	0.606
	Huitong	A	$y=21.2e^{-0.0290x}$	0.92	<0.001	$y=2.47 \times 10^{-4}x^2-0.0324x+1.59$	0.93	<0.001	$y=4.76 \times 10^{-5}x^2-0.0046x+0.178$	0.40	<0.001
		B	$y=155x^{-0.934}$	0.97	<0.001	$y=4.27x^{-0.425}$	0.97	<0.001	$y=0.216x+0.176$	0.27	0.228
Grassland	Miaoxian	A	$y=197x^{-0.905}$	0.96	0.001	$y=11.3x^{-0.734}$	0.95	0.001	$y=8.15 \times 10^{-5}x^2-0.0094x+0.409$	0.96	0.040
		B	$y=154x^{-0.884}$	0.94	0.001	$y=13.4x+0.187$	0.99	<0.001	$y=6.93 \times 10^{-5}x^2-0.0088x+0.371$	0.89	0.112
	Shennongjia	A	$y=77.3x^{-0.503}$	0.89	0.004	$y=4.09x^{-0.393}$	0.84	0.011	$y=4.32x^{-0.0531}$	0.51	0.176
		B	$y=64.9e^{-0.0144x}$	0.84	0.010	$y=4.69e^{-0.0126x}$	0.91	0.003	$y=e^{-0.409/x+1.57}$	0.24	0.407
	Xishuanbanna	A	$y=39.5x^{-0.442}$	0.87	<0.001	$y=2.99x^{-0.365}$	0.84	<0.001	$y=1.15x+0.181$	0.90	<0.001
		B	$y=47.4x^{-0.345}$	0.64	0.017	$y=7.11x^{-0.520}$	0.88	0.001	$y=9.13 \times 10^{-5}x^2-0.0117x+0.734$	0.89	0.004
	Changbai	A	$y=1591x^{-1.53}$	0.77	0.004	$y=77.4x^{-1.23}$	0.77	0.004	$y=4.00 \times 10^{-4}x^2-0.0516x+1.81$	0.59	0.264
		B	$y=305x^{-1.13}$	0.92	0.001	$y=13.8x^{-0.855}$	0.91	0.001	$y=6.04x+0.121$	0.85	0.026
	Haibei	A	$y=70.9e^{-0.0348x}$	0.80	0.003	$y=5.96e^{-0.0314x}$	0.91	<0.001	$y=-0.0536 \ln x+0.996$	0.31	0.149
		B	$y=122x^{-0.364}$	0.80	0.003	$y=11.1x^{-0.390}$	0.92	<0.001	$y=4.89 \times 10^{-6}x^2-0.0031x+1.03$	0.68	0.059
	Inner Mongolia	A	$y=37.1x^{-0.419}$	0.82	<0.001	$y=4.45x^{-0.409}$	0.63	0.002	$y=2.39 \times 10^{-5}x^2-0.0037x+0.324$	0.99	0.001
		B	$y=20.8e^{-0.0239x}$	0.99	<0.001	$y=-2.66 \times 10^{-4}x^2-0.0374x+1.81$	0.98	0.023	$y=-6.06 \times 10^{-6}x^2-0.0012x+0.295$	0.98	0.017
Desert	Sanjiang	A	$y=789e^{-0.102x}$	0.77	0.002	$y=33.2e^{-0.0723x}$	0.62	0.012	$y=-3.82 \times 10^{-4}x^2+0.0036x+0.961$	0.64	0.048
		B	$y=183e^{-0.0879x}$	0.75	0.006	$y=12.7e^{-0.0593x}$	0.77	0.004	$y=6.19 \times 10^{-4}x^2-0.0451x+1.69$	0.59	0.111
	Cele	A	$y=1.34e^{0.00183x}$	0.66	0.049	$y=4.24 \times 10^{-6}x^2+3.48 \times 10^{-4}x+0.183$	0.80	0.090	$y=-1.42 \times 10^{-3}x^2+7.76 \times 10^{-4}x+0.563$	0.67	0.190
		B	$y=1.49e^{-0.00134x}$	0.37	0.199	$y=-8.55 \times 10^{-6}x^2+8.18 \times 10^{-4}x+0.179$	0.22	0.691	$y=-1.05 \times 10^{-3}x^2-0.0011x+0.561$	0.84	0.061
	Erdos	A	$y=6.06x+0.706$	0.99	0.001	$y=0.293x+0.122$	0.74	0.060	$y=-6.96 \times 10^{-6}x^2-0.0011x+0.244$	0.95	0.048
		B	$y=-0.638 \ln(x)+3.11$	0.94	0.006	$y=-0.0693 \ln x+0.354$	0.96	0.003	$y=2.98 \times 10^{-5}x^2-0.0019x+0.245$	0.59	0.408
	Fukang	A	$y=4.56x^{-0.427}$	0.94	0.007	$y=2.03 \times 10^{-5}x^2-0.00297x+0.235$	0.90	0.098	$y=6.56 \times 10^{-5}x^2-0.0080x+0.476$	0.95	0.046
		A	$y=e^{2.72/x+0.732}$	0.93	0.008	$y=0.991/x+0.243$	0.96	0.004	$y=6.20 \times 10^{-6}x^2+5.57 \times 10^{-4}x+0.260$	0.67	0.326
	Naiman	B	$y=2.98x+0.414$	0.96	0.003	$y=-0.0235 \ln x+0.176$	0.98	0.001	$y=4.01 \times 10^{-5}x^2-0.0053x+0.249$	0.90	0.103
		A	$y=12.6x^{-0.937}$	0.81	0.003	$y=-0.0980 \ln x+0.466$	0.86	0.001	$y=7.82 \times 10^{-6}x^2-0.0015x+0.159$	0.97	0.027
	Shapotou	B	$y=4.23x+0.240$	0.97	<0.001	$y=-0.0406 \ln x+0.233$	0.90	0.001	$y=0.200x+0.104$	0.96	0.003
		A	$y=5.64 \times 10^{-5}x^2-0.0113x+1.13$	0.92	0.082	$y=-0.0241 \ln x+0.219$	0.92	0.010	$y=0.341/x+0.278$	0.99	0.001

Notes: A and B represent different plots with different dominant species or ecosystem types in the same ecological station; y and x represent the content of soil organic carbon (g/kg) and soil depth (cm), respectively



N, and P in all plots was well fitted by different curve functions, although the fitting functions differed among different plots. The SOC profile distribution was better fitted with exponential functions in the artificial agricultural ecosystems only excluding the plots of Changwu and Naiman stations from 38 plots. Among the type of the best fitting functions, there were 53% of the exponential function and 45% for the power function. Similarly, soil TN can be fitted by quadratic functions in all plots ( $P < 0.05$  for all plots), with 76%, 11%, and 8% of the best fitting functions being respectively quadratic, exponential and logarithmic functions. In total, the power functions fitted the SOC profile distributions well in 36 natural ecosystems plots (desert, forest, and grassland) except for the plots of Cele station ( $P < 0.05$  for all plots). The best fitting functions were about 44% for the power function, 36% for the exponential function, and 6% for the S function and 8% for inverse functions. Soil TN content was best fitted by power functions in the natural ecosystems, with the best fitting functions being 31%, 19%, and 19% for power, exponential, and quadratic functions, respectively.

Quadratic functions well depicted the profile distribution of TP in the 0–100 cm soil for all 74 plots (Table 2, Fig. 4). The best fitting functions for TP were the quadratic function (97%) followed by logarithmic functions in the artificial ecosystems, and the quadratic functions (67%) followed by the inverse (19%) and S (6%) functions in the natural ecosystems.

Overall, the best fit functions for soil C and N content clearly differed between artificial (agriculture) and natural (desert, forest, and grassland) ecosystems; however, the vertical distribution functions of soil P content were similar in both artificial and natural ecosystems.

### 3.3 Stoichiometric ratios of C, N, and P in soil

Stoichiometric ratios of C, N, and P of soils are ordered as follows: grassland > forest > agriculture > desert (Table 3). The similar vertical distributions of soil C, N, P, and their best fitting functions in both artificial and natural ecosystems indicated the close coupling relationships between soil C, N, and P at a large scale. The stoichiometric ratios of C, N, and P were also associated with climate to some extent. The ratios generally were lower in the warm temperate and north subtropical zones, higher in the temperate and south subtropical zone, and highest in the Tibetan Plateau (Table 3).

**Table 3** Stoichiometric ratios of soil C, N, and P in Chinese terrestrial ecosystems

Classification	Type	C : N	N : P	C : P
Ecosystem type	Agriculture	9	1	13
	Forest	13	3	41
	Grassland	14	6	83
	Desert	8	1	5
Climatic zone	Temperate	13	3	40
	Warm temperate	8	1	10
	North subtropical	11	1	12
	Mid-subtropical	10	2	21
	South subtropical	14	3	42
	Tibetan Plateau	13	4	54

Note: Tibetan Plateau have been classified individually due to especial climate and altitude

## 4 Discussion

### 4.1 Fitting functions for profile distribution of soil C, N, and P

This study established a set of the fitting functions for the profile distribution of C, N, and P in the 0–100 cm soil layer in the Chinese terrestrial ecosystems. The exponential function, quadratic function, and quadratic function were found to be more suitable to characterize the vertical distribution of soil C, N, and P, even though the best fitting functions differed among different ecosystems. Many studies have demonstrated that the soil C, N, and P content decreases with increasing soil depth (Panda *et al.*, 1988; Bowman and Savory, 1992; Hobley *et al.*, 2013). Furthermore, Zinke (1978) developed the cumulative log-log models to describe SOC profiles. Jobbagy and Jackson (2000) used the logarithm functions to construct the relationships between SOC and soil depth using three global databases. Our study suggests that it should be cautious to use a general function to simulate the change of SOC content with increasing soil depth because the best fitting functions differed among ecosystems.

### 4.2 Differences between artificial and natural ecosystems

The fitting functions for the profile distribution of soil C, N, and P content in the 0–100 cm appeared to differ in artificial and natural ecosystems. The exponential and quadratic functions fitted the C, N, and P profile distribution well in agricultural soils. In contrast, the power and quadratic functions were the best fitting

functions for natural soils (forest, grassland, and desert soils). Turner (1989) reported that the soil environment may alter as a result of changes in natural and ecological processes under different ecosystem types. In the artificial ecosystems, natural and human factors jointly influence the distribution of soil nutrients (Wang *et al.*, 1996), with fertilization and cultivation exerting stronger impacts on the upper soil (Coleman *et al.*, 1993; Drinkwater and Snapp, 2007). Furthermore, Crow *et al.* (2009) reported that the C and P content in the surface soil was directly related to the input of litter and mulch in farmlands. In the deep soil layer, soil genesis becomes more important. In natural ecosystems, soil forming factors, such as climate, topography, parent material, biology, and time are important for soil C, N, and P vertical distributions (Sinsabaugh *et al.*, 1993; Olander and Vitousek, 2000; Xia *et al.*, 2014). Furthermore, soil P content is regulated by various biological factors and disturbance in the short-term, while geochemical factors of the parent rock represent the main controlling factors in the long-term (Kellogg and Bridgham, 2003).

#### 4.3 Coupling relationship of soil C, N, and P vertical distribution

The best fitting functions observed here imply the close coupling relationships among the C, N, and P storage in the 0–100 cm soil profile. Ecological stoichiometry (or soil C : N : P ratios) are expected to provide a new approach to analyze the coupling cycles of soil C, N, and P (Sterner and Elser, 2002). Our findings showed that the C : N, N : P and C : P ratios ranked in order from grassland > forest > agriculture > desert. The observed differences were mainly due to differences in the element ratios of vegetation consumption and release from soil and atmosphere (Sterner and Elser, 2002).

Some studies have suggested that climate factors play an important role in the vertical distribution of soil C, N, and P by controlling soil development, biota, and their interactions (Chadwick *et al.*, 1999; Oleksyn *et al.*, 2003; Vitousek, 2004). The stoichiometric ratios of C, N, and P were associated with climate in this study. The general trends were lower in warm temperate and north subtropical zones, and higher in temperate (or Tibetan Plateau) and south subtropical zones. The tropical and subtropical ecosystems with high productivity do not maintain relatively high soil C content due

to the stronger effects of temperature on SOM decomposition, but high temperatures and ample precipitation in these systems may result in high P leaching rates and P occlusion of highly weathered soils (Vitousek *et al.*, 1987; Neufeldt *et al.*, 2000; Zhang *et al.*, 2005), leading to the highest C : P and N : P ratios. In contrast, the dry and cool climate of temperate deserts results in low productivity, lower soil C and N content, and low P loss through leaching, along with higher soil P content. Consequently, temperate deserts have the lowest soil C : P and N : P ratios among all climatic zones (Tian *et al.*, 2010). In addition, Walker and Adams (1958) reported that C : P ratios declined much faster than the C : N ratios with increasing soil depth. One possible explanation for this phenomenon is that soil P mainly derives from soil weathering, in addition to soil available P migrating through plants and surface accumulation.

The delineation of fitting functions for the vertical distribution of soil C, N, and P in typical terrestrial ecosystems is of great significance. First, the functions provide a scientific basis and new approaches to accurately estimate the storage of soil C, N, and P at regional and national scales. These fitting functions are of significance for scientists to integrate sparse data from different sources with different soil depths to estimate C storage at large scale. Second, the coupling cycles of C, N, and P have been considered as one of the hot spots of global climate change study (Lal, 2004; Schipper *et al.*, 2004).

## 5 Conclusions

Regression functions were built for the profile distribution of soil C, N, and P content in the 0–100 cm depth in the Chinese terrestrial ecosystems. The best fitting functions with respect to soil C, N, and P content differed, to some extent, in different plots, especially between artificial and natural ecosystems. In both the artificial and natural ecosystems, the best fitting functions were exponential and power functions for C, quadratic and power functions for N, and quadratic functions for P. The best fitting functions for soil C, N, and P content indicated a close coupling relationship among the storage and turnover of soil C, N, and P. The stoichiometric ratios of soil C, N, and P content ranked as follows: grassland > forest > agriculture > desert, and are associated with climate. This study is the first to build the

fitting functions of vertical profiles at the national scale. The study provides an important tool to accurately estimate soil C, N, and P storage at a large scale, and give new insights into the coupling cycles of soil C, N, and P.

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